

Lessons Learned from the Assimilation of Cloud-affected SEVIRI Radiance Observations at NCEP

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Funding: NOAA GOES-R visiting scientist project

Background

- Large numbers of satellite radiances contain cloud and precipitation signal
- Most data assimilation system in current NWP screen the radiance data and reject cloud/rain contaminated observations, only assimilate the clear-sky radiance
- If cloudy radiances can be properly used, potentially will improve the forecasts of temperature, wind, moisture and cloud fields
- NCEP has been made using cloudy AMSU-A radiance data in operational
- Starting work with IR cloudy radiance data assimilation for IASI, CrIS, SEVIRI

Assimilate cloud-affected IR radiance

Simplified system:

very simple cloud representation
currently limited to overcast scenes
no information on clouds taken from model
no back interaction with model via physics



Advanced system:

very complex cloud representation
all cloud conditions treated
information on clouds taken from model
back interaction with model via physics

from: McNally,2012







- For assimilating overcast radiance, observation-centered cloud top height is used as a control variable.
- ✤ In addition to modified $\partial T_B / \partial T(p)$, $\partial T_B / \partial q_v(p)$, etc., the minimization now incorporates the CTP Jacobian, $\partial T_B / \partial p_{cld}$.
- Under the grey-body assumption, the partially cloudy observation can then be considered for a single, fractional cloud as:

 $R_i = R_i^{clr} (1 - N) + R_i^{overcast} N$

The initial cloud top height and cloud fraction can be obtained from the GSI cloud detection algorithm, and will be adjusted during the minimization.



- Overcast conditions are least ambiguous in the radiance data
- Cloud control vector simplified to a single number (cloud top pressure), instead of complex hydrometer variables
- Do not need to consider cloud overlap situation
- Termination of Jacobians at cloud top provides new information, and measure temperature above clouds better than in clear sky
- No cross-talk between cloud and surface skin sink variables

Example of Analysis Increment





IR cloud detection

• The simple RTM calculates a radiance at channel *I under the single-layer cloud assumption* as:

$$R_i = R_i^{clr} (1 - N) + R_i^{overcast} N$$

• Using the minimum residual method (Eyre and Menzel, 1989), minimize the cost function

$$J = \sum_{i} (R_i^{obs} - R_i)^2$$

• By setting
$$\frac{dJ}{dN} = 0$$
, $N = \frac{\sum_{i} (R_i^{clr} - R_i^{obs})(R_i^{clr} - R_i^{overcast})}{\sum_{i} (R_i^{clr} - R_i^{obs})}$

• Cloud top height was determined at a discrete model level where $R_i^{overcast}$ was calculate

O-B of 10.8 μm band BT.

10.000

6 667

Before Bias Correction



O-B no/bc (K) Data Count= 5289

After Bias Correction



O-B w/bc (K) Data Count=

0.000

-3.333

5289

6.667

10.000

3.333

O-B Histogram

Model has warm bias/observation has cold bias



O-B of 10.8 μm band BT.

Before Bias Correction



O-B no/bc (K) Data Count= 5289

After Bias Correction



O-B w/bc (K) Data Count=

0.000

3 3 3 3

3 333

5289

Model warm bias
 would lead to higher
 estimates of cloud top

Cloud top pressure



Cloud Top Pressure (mb)

100.000

233 333



CLDTP seveld_m10

Cloud Top Pressure (mb)

900 000

 366.667
 500.000
 633.333
 766.667
 900.000
 100.000
 233.333
 366.667
 500.000
 633.333
 766.667

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Cloud Fraction



- Warm model bias would lead to smaller cloud fraction than really exist.
- This might cause the sharp temperature Jacobian at the wrong height and with wrong amplitude.
- So correct as many bias in models and measurements as possible before estimating the cloud top height and cloud fraction

Why exploring SEVIRI radiance?

- EUMESAT Meteosat Second Generation (MSG) Geostationary Satellite SEVIRI has the similar channels, resolutions as GOSE-R ABI
- So hopefully the lessons and experiences we learned from SEVIRI can be used for assimilating cloud-affected GOES-R ABI radiance into NCEP operational data assimilation system

SEVIRI vs GOES-R ABI

ABI chn #	SEVIRI chn #	Wave length (µm)) Main observation application		
1		VISo.47	Aerosol detection , visibility estimation		
2	1	VISo.6	Surface, clouds, wind fields		
3	2	VISo.8	Surface, clouds, wind fields		
4		NIR1.4	Very thin cirrus clouds		
5	3	NIR1.6	Surface, cloud phase		
6		IR2.3	Aerosol, cloud particle size		
7	4	IR3.9	Surface, clouds, wind fields]	
8	5	WV6.2	Water vapor, high level clouds, atmospheric instability		
9		WV7.0	Water vapor, atmospheric instability	1	
10	6	WV _{7.4}	Water vapor, atmospheric instability		
11	7	IR8.5	Surface, clouds, atmospheric instability		
12	8	IR9.7	Ozone		
13		IR10.35	Low-level moisture,cloud particle size		
14	9	IR11.2	Surface, clouds, wind fields, atmospheric instability		
15	10	IR12.3	Surface, clouds, atmospheric instability		
16	11	IR13.3	Cirrus cloud height, atmospheric instability		

SEVIRI ~3 km footprint GOES-R ~2 km footprint

2 water vapor - channels are assimilated in GDAS

Forecast and data assimilation system

Hourly-Updated NAM Forecast System for Africa

- o NAMRR: NAM Rapid Refresh
- Hourly updates
- Cycling 12 km parent and 4 km nest
- Hybrid ensemble-3DVar via Global Data Assimilation System's EnKF members

NAMRR





EUMETSAT SEVIRI ASR Product



Assimilated SEVIRI Brightness Temperatures

Water Vapor Band (6.2 µm)

<u>Clear-sky BT</u>

Overcast cloudy BT



Cloudy observations add much more data over Atlantic Ocean A few more data assimilated in mid-Africa

Analysis difference w/o overcast cloud

<u>700mb Temperature Difference (m/s)</u>



Temperature difference increased with the catch-up time



300mb Specific Humidity Difference (g/kg)



Humidity difference for mid-Africa was observed at the last cycle.

Model Forecast Validation

- Assess NAMRR forecast for accuracy using SEVIRI infrared brightness temperatures computed from model fields using CRTM
- This allow a direct comparison of forecasts with satellite data, especially for the data sparse area
- Primary focus is to assess the accuracy of cloud and moisture fields
- Using EUMETSAT SEVIRI All-Sky-Radiance (ASR) product from the 6.2 and 7.3 μm (water vapor sensitive) and 10.8 μm window channel (clouds and surface temperatures) and 13.4 μm (cirrus cloud)
- In addition of the model forecast temperature and specific humidity profiles, cloud liquid water, cloud ice water, rain water and snow/graupel profiles are used by CRTM to compute the brightness temperature hourly



268 200 287.650 248 750 233 200 246 833 260 467 249 987 261 400 192 300 219 567

4km Model hourly forecast brightness temperature for window channel



The location of cloud from 4km model forecast is consistent with observations for both clear and cloudy experiments. But model forecast for some convective clouds is very weak, except for Lake Victoria storm. There is no big difference Between 'clear' and 'cloudy'.

Model Forecast RMSE

$$\text{RMSE} = \sqrt{\frac{1}{N} \sum_{i=1}^{n} (F_i - O_i)^2}$$

Calculated with the entire 4km nest o-6 hs model forecast and SEVIRI all-sky radiance



Experiment with overcast cloudy radiance has the smaller forecast RMSE for 1-6 hours.

Conclusions

- 1. Through adjusting the cloud top height, overcast cloudy SEVIRI radiance are assimilated in both global and regional forecast, and has expected impact.
- 2. Cloud detection scheme is very important to assimilate IR overcast cloud radiance. The cloud top height is the crucial element to ensure the model simulated overcast radiance more accurate
- 3. Assimilation only overcast cloud IR radiance is not enough to improve the storm forecast for less data
- 4. NCEP NAMRR model forecast has the systematically weak convective clouds for Africa region with current physics scheme(Ferrier et al. (2002, 2011))



 Perform the direct assimilation of all-sky SEVIRI radiance with the hydrometeor variables (cloud ice, cloud water, rain water, snow/grauple...)

Advanced system:

very complex cloud representation
all cloud conditions treated
information on clouds taken from model
back interaction with model via physics



7.3 µm All-sky Brightness Temperature



Brightnes	ss reinperature (K)		Count-	JOTT AVG BT-		200.1
203.600	214.717	225.833	236.950	248.067	259.183	270.3



Brightne	rightness Temperature(K) Count=			5611 AVG BT=		253.3	
203.600	214.717	225.833	236.950	248.067	259.183	270.3	





